

## Introduction and Motivation

- Previous work suggested that mare basin-related extension on the Moon largely ended ~3.6 Ga and contractional deformation ended ~1.2 Ga<sup>1-4</sup>
- Wrinkle ridges are often associated with mascons (large positive gravity anomalies)<sup>2,3,5</sup> yet ridges occur in Mare Frigoris even where a large mascon is not observed
- Lunar Reconnaissance Orbiter Camera (LROC) enables the discovery of tectonic landforms at scales not previously imaged<sup>6,7</sup>
- Landform morphology and stratigraphic relationships imply a complex history of deformation of the Moon

## Landform Classifications

- Lobate Scarp: A simple curvilinear, asymmetric hill formed by near-surface fault<sup>5,7-10</sup> (Fig. 1a)
- Wrinkle Ridge: A complex curvilinear, asymmetric hills formed by folding over a blind fault or faults<sup>2,3,5,11-15</sup> (Fig. 1b)
- Graben: A trough formed between two normal faults<sup>1,6,11</sup> (Fig. 1c)

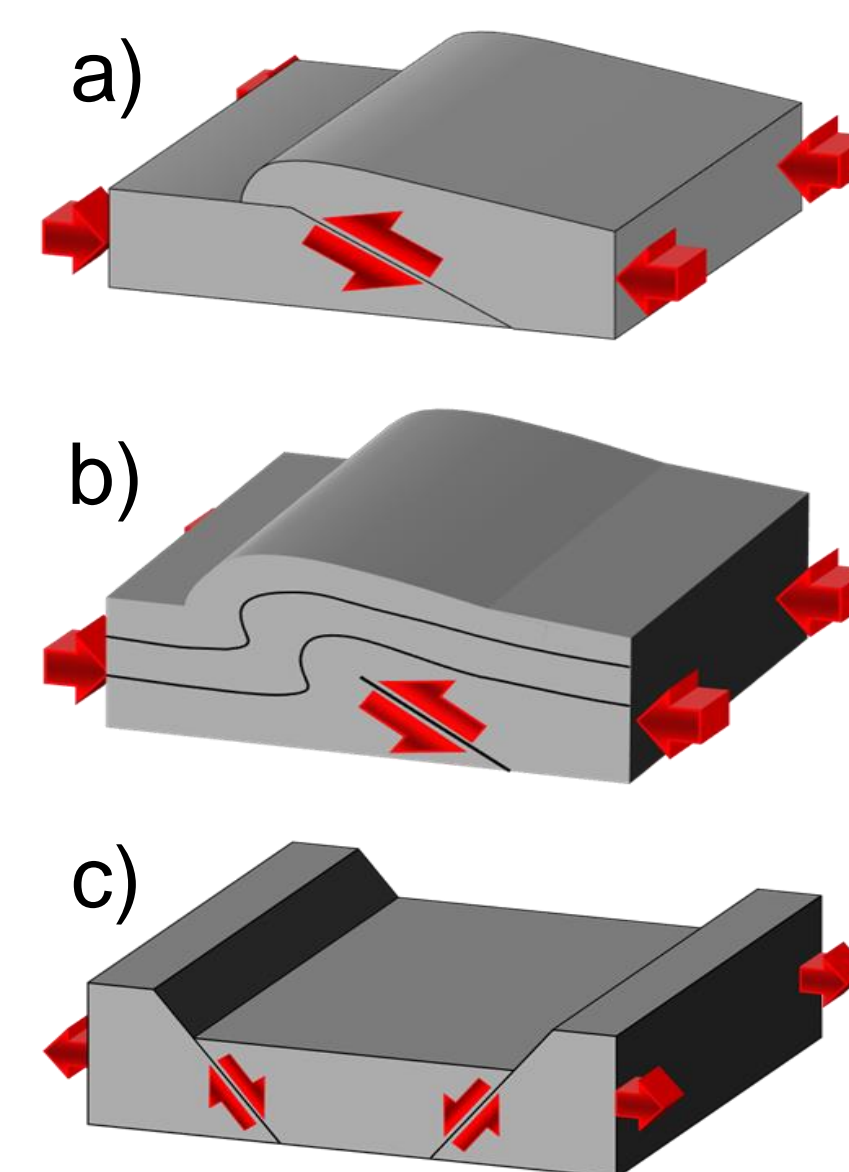


Fig. 1: Block diagrams of a) lobate scarp, b) wrinkle ridge, and c) graben

## Data and Methods

- LROC Narrow Angle Camera (NAC) images with meter-scale resolution<sup>16</sup>
- Nearly continuous NAC image coverage from 45°N to 65°N and 45°W to 45°E
- Map tectonic landforms using ArcGIS
- Correlate landform distribution with GRAIL free air gravity anomaly<sup>17</sup>
- Find and measure small crosscut craters and classify degradation state to determine age<sup>18</sup>

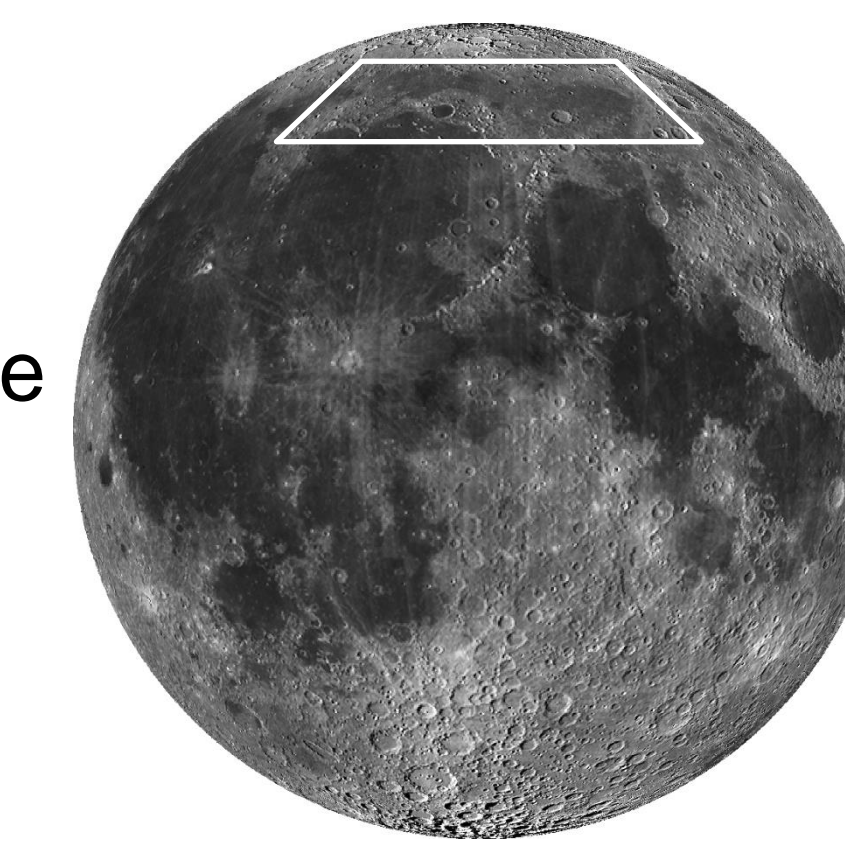


Fig. 2: LROC WAC global context<sup>19</sup> of Mare Frigoris (trapezoid)

## Landforms in Western Mare Frigoris

- Parallel set of wrinkle ridges that trend NW/SE (black lines)
- These ridges are overprinted by numerous large craters (up to hundreds of meters in diameter) and have broadly undulating slopes, indicating a relatively old age
- Likely formed soon after mare basalt emplacement (2.6-3.8 Ga with most between 3.4-3.8 Ga<sup>20,21</sup>)
- Parallel ridges are not consistent with mascon-induced flexure<sup>2,3</sup> and the negative or very small positive gravity anomaly<sup>17</sup>, nor do the askew orientations suggest an Imbrium outer ring collapse<sup>20</sup>
- 2 other clusters of tectonic landforms (white circles) occur independent of the parallel series of ridges
- These wrinkle ridges and lobate scarps tend to be smaller, often only tens of meters in relief, and have sharp changes in slope, suggesting a relatively young age
- Many small graben found in the back limbs of these crisp ridges/scarps

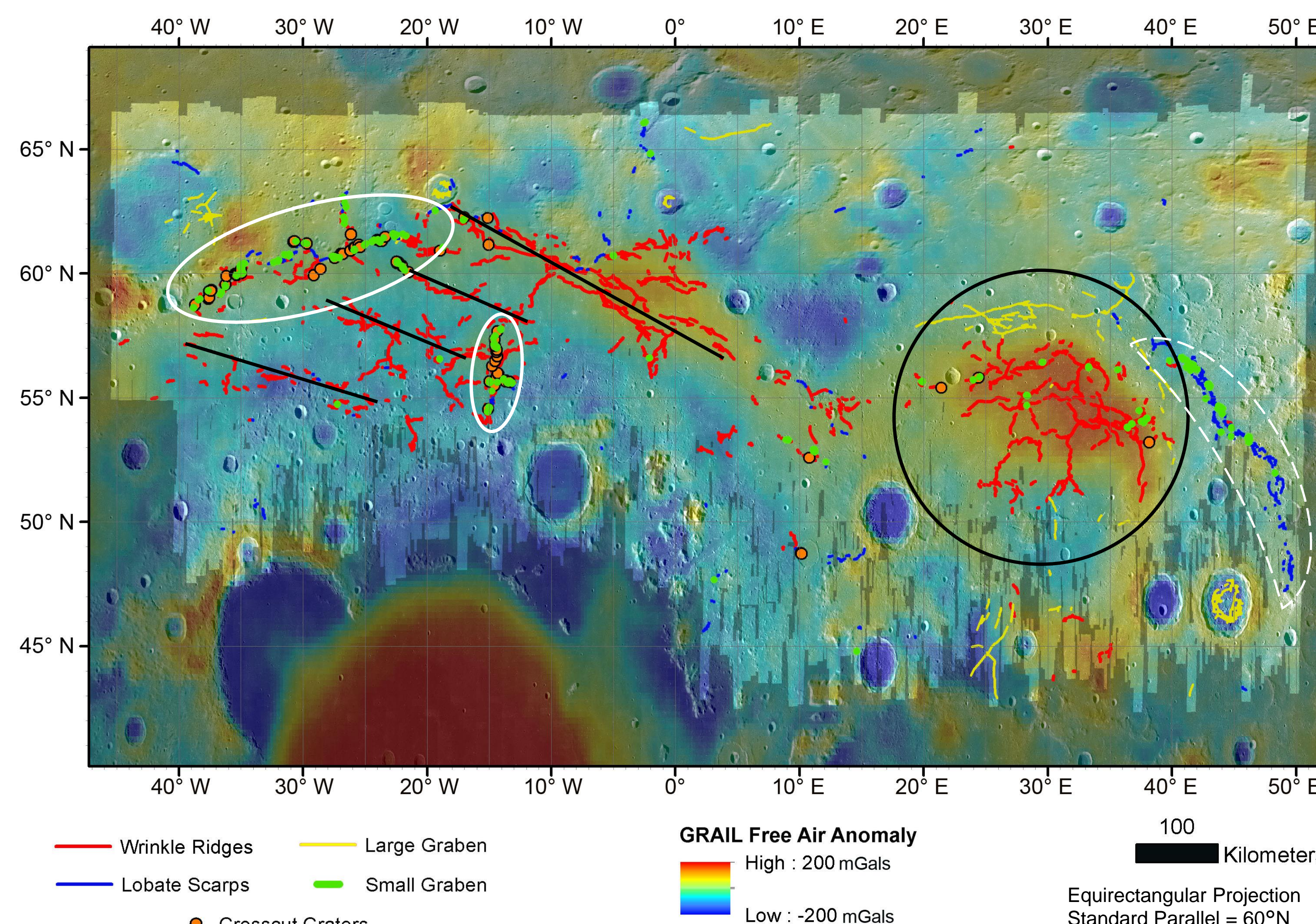


Fig. 3: Tectonic map of Mare Frigoris over GRAIL Free Air Gravity Anomaly<sup>17</sup> and LROC WAC shaded relief<sup>19</sup>

## Landforms in Eastern Mare Frigoris

- Polygonal pattern of wrinkle ridges bounded on the north and east sides by large arcuate graben (black circle)
- These ridges and large graben are overprinted by numerous large craters (up to hundreds of meters in diameter) and have broadly undulating slopes, indicating a relatively old age
- Likely formed soon after mare basalt emplacement (2.6-3.8 Ga with most between 3.4-3.8 Ga<sup>20,21</sup>)
- Consistent with the traditional mascon-induced flexure<sup>2,3</sup> and correlated with positive gravity anomaly<sup>17</sup>
- Two ~250 km sub-parallel series of an echelon, lobate scarps (dashed white wedge) extends into the highlands from eastern Mare Frigoris, with numerous small-scale graben in the back limbs of scarps
- Like other lobate scarps globally, this series is inferred to have formed within the last 1.0 Ga as a result of a global compressional stress from cooling and radial contraction of the Moon's interior<sup>5,7,22-24</sup>

## Absolute Age Constraints from Small Crosscut Craters

- Small craters quickly destroyed from impact gardening
- Craters ≤80 m in diameter are ≤1.0 Ga<sup>18</sup>
- Wrinkle ridges observed crosscutting at least 70 craters 21-100 m in diameter, a few still showing bright ejecta
- Calibrated degradation rates for small craters<sup>18</sup> suggest observed crosscut craters as young as ~40 Ma (±3x)
- Seismic shaking would decrease retention age, so ages are likely overestimated

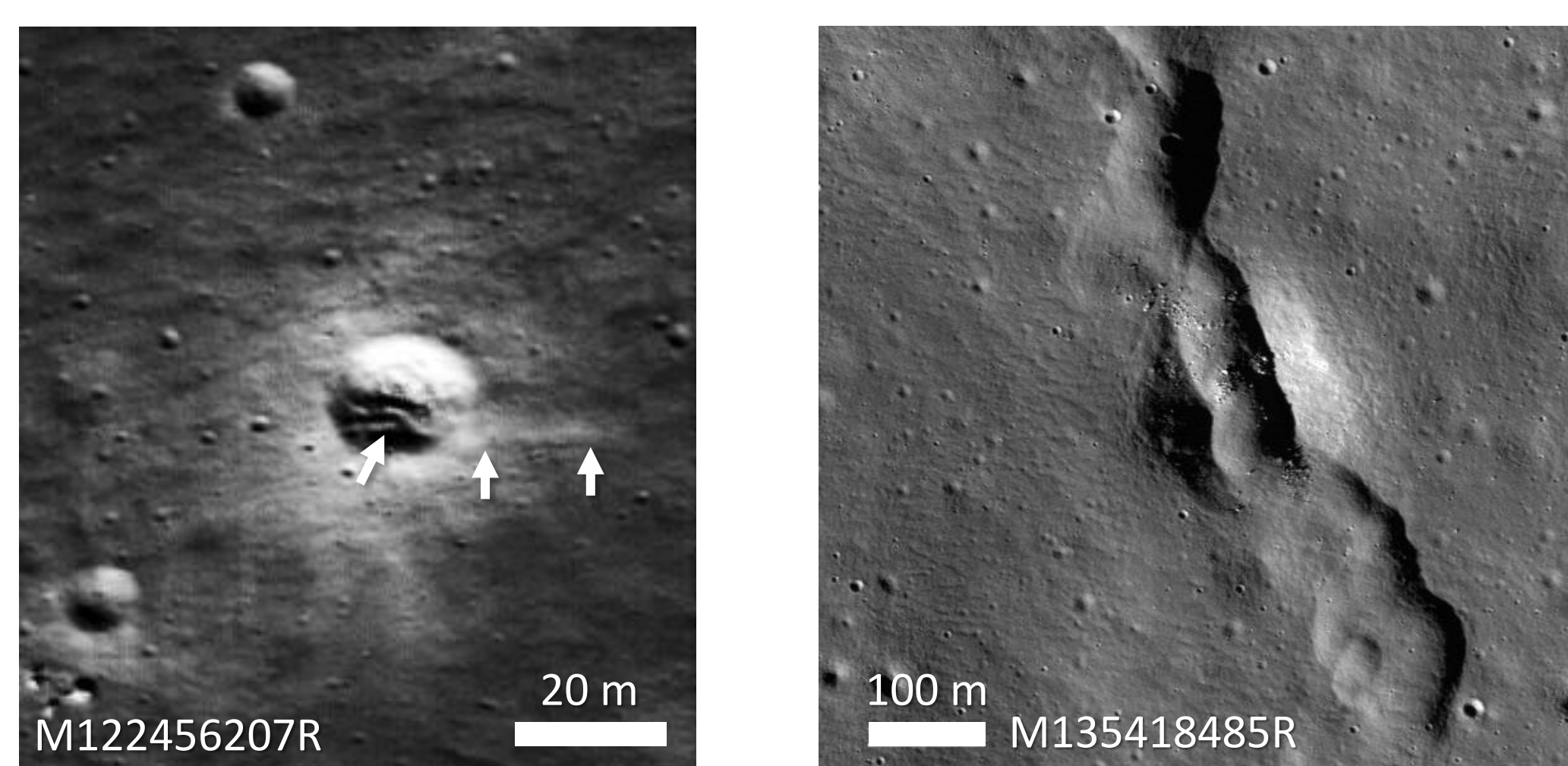


Fig. 4: LROC NAC images of small craters crosscut by wrinkle ridges

## Constraints from Small Graben

- Meter-scale graben occur in back limbs of many morphologically crisp ridges/scarps, and usually either parallel or perpendicular to the associated ridge/scarp
- Parallel graben consistent with flexural bending in back limb<sup>6,7,11</sup>
- Perpendicular graben akin to tension gashes<sup>11</sup>
- A few have pit crater chains similar to Vitello graben<sup>6</sup>
- Based on regolith infill rates for shallow troughs, meter-scale graben estimated to be <50 Ma<sup>6,25</sup>
- Suggests associated ridges/scarps active within <50 Ma

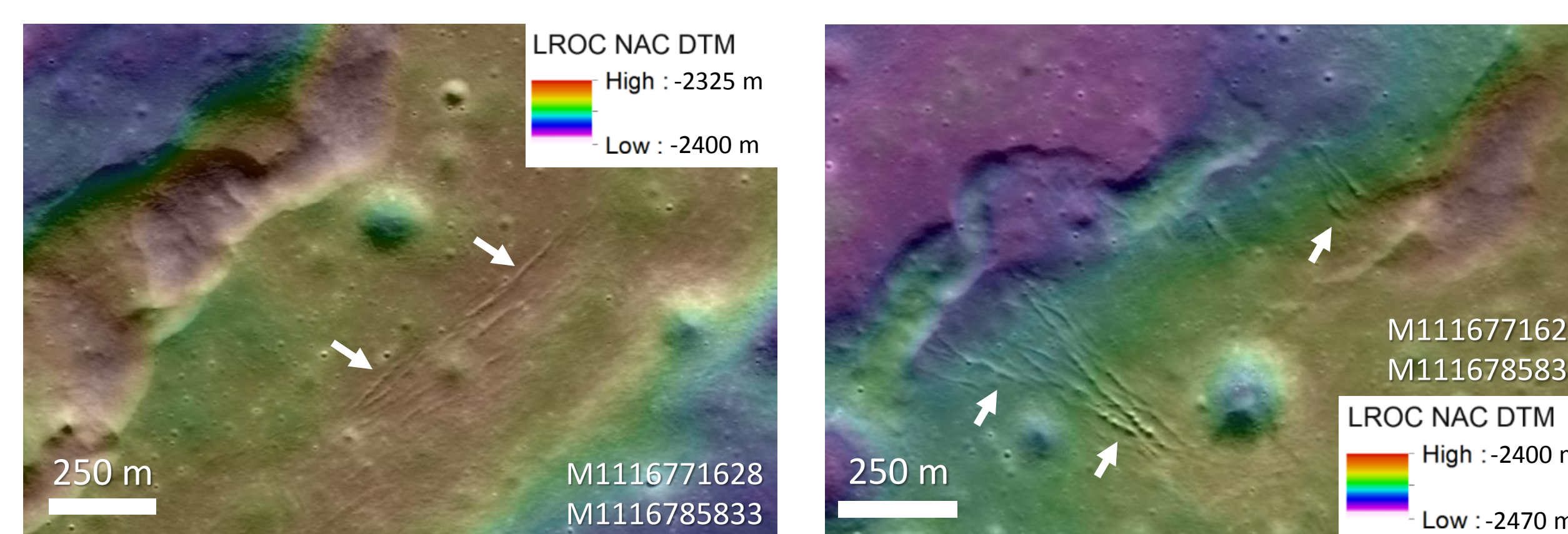


Fig. 5: LROC NAC digital terrain models of small graben (arrows), parallel (left) and perpendicular (right) to their associated wrinkle ridges

## Conclusions

- Mare Frigoris is a tectonically diverse region with a complex deformational history
- Older wrinkle ridges and large graben in eastern Mare Frigoris likely formed due to mascon flexure and subsidence
- Older wrinkle ridges in western Mare Frigoris not consistent with mascon flexure and subsidence
- Lobate scarps and perhaps some wrinkle ridges may be due to global contractional stress, were active within the past 1.0 Ga, and might still be active today

## References

- <sup>1</sup>Lucchitta B. K. and Watkins J. A. (1978) LPS 9, 3459-3472. <sup>2</sup>Solomon S. C. and Head J. W. (1979) JGR 84, 1667-1682. <sup>3</sup>Solomon S. C. and Head J. W. (1980) Rev. Geophys. & Space Phys. 18, 107-141. <sup>4</sup>Hiesinger H. et al. (2003) JGR 108, E001985. <sup>5</sup>Watters T. R. and Johnson C. L. (2010) in Planetary Tectonics, Cambridge Univ. Press, 121-182. <sup>6</sup>Watters T.R. et al. (2012) Nature Geosci., doi:10.1038/ngeo1387. <sup>7</sup>Watters T.R. et al. (2010) Science, 329, 936-940. <sup>8</sup>Binder A.B. and Gunga H.C. (1985) Icarus, 63, 421-441. <sup>9</sup>Banks M.E. et al. (2012) JGR 117, doi:10.1029/2011JE003907. <sup>10</sup>Williams N.R. et al. (2013) JGR 118, 224-233. <sup>11</sup>Plescia J. B. and M. P. Golombek (1986) GSA Bull. 97, 1289-1299. <sup>12</sup>Watters T. R. (1988) JGR, 93, 10236-10254. <sup>13</sup>Golombek M. P. et al. (1991) LPS 21, 679-693. <sup>14</sup>Schultz R. A. (2000) JGR 105, 12035-12052. <sup>15</sup>Watters T. R. (2004) Icarus 171, 284-294. <sup>16</sup>Robinson M. S. et al. (2010) Space Sci. Rev. 150, 81-124. <sup>17</sup>Zuber M. T. et al. (2013) Science 339, 668-671. <sup>18</sup>Moore H. J. et al. (1980) Moon and Planets 23, 231-252. <sup>19</sup>Speyerer E. et al. (2011) LPSC 42, 2387. <sup>20</sup>Whitford-Stark J. L. (1990) LPSC 20, 175-185. <sup>21</sup>Hiesinger H. et al. (2010) JGR 115, E03003. <sup>22</sup>Solomon S. C. and Chaiken J. (1976) LPS 7, 3229-3243. <sup>23</sup>Kirk R. L. and Stevenson D. J. (1987) JGR 94, 12133-12144. <sup>24</sup>Watters T. R. et al. (2013) LPSC 45, this volume. <sup>25</sup>Arvidson, R. et al. (1975) Moon 13, 67-79.

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